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Texture development and elastic stresses in magnesiowüstite at high pressure

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Abstract Cubic magnesiowüstite has been deformed in a diamond anvil cell at room temperature. We present results for $(\text{Mg}_{0.4}\text{Fe}_{0.6})\text{O}$, $(\text{Mg}_{0.25}\text{Fe}_{0.75})\text{O}$, and $(\text{Mg}_{0.1}\text{Fe}_{0.9})\text{O}$ up to 37, 16, and 18 GPa, respectively. The diffraction images, obtained with the radial diffraction technique, are analyzed using both single peak intensities and a Rietveld method. For all samples, we observe a [100] fiber texture but the texture strength decreases with increasing iron content. This texture pattern is consistent with $\{110\}\langle 1-10 \rangle$ slip. The images were also analyzed for stress, elastic strains, and elastic anisotropy. In general, the stress measured in magnesiowüstite samples is lower than previously measured on MgO. The elastic anisotropy deduced from the X-ray measurements shows a broad agreement with models based on measurements with other techniques.

Keywords (Mg,Fe)O · Texture · DAC · Synchrotron X-rays · Deformation mechanisms · Magnesiowüstite · Elasticity

Introduction

Magnesiowüstite [(Mg,Fe)O] is a major component of the lower mantle and its deformation behavior is important for understanding the rheology of the deep earth. The amount of Fe present in (Mg,Fe)O in the lower mantle has been estimated to be 10–45 mol% FeO, based on high P – T element partitioning experiments (e.g. Mao et al. 1997; Andrault 2001 and

references therein). Both wüstite, which has a nonstoichiometric formula Fe_{1-x}O (e.g. McCammon 1993), and periclase (MgO) have a NaCl (B1) structure at ambient conditions. Intermediate and iron-rich compositions of magnesiowüstite show a phase transition to a rhombohedral phase at high pressure (Yagi et al. 1985; Richet et al. 1989; Kondo et al. 2004; Lin et al. 2003). There is also a pressure-induced electronic spin transitions of iron in (Mg,Fe)O at high pressure (Sherman 1988, 1991; Sherman and Jansen 1995; Cohen et al. 1997; Badro et al. 2003; Lin et al. 2005; Speziale et al. 2005), and a possible dissociation of intermediate compositions into magnesium-rich and iron-rich components (Dubrovinsky et al. 2000, 2001, 2005). Magnesiowüstite is of central importance for geodynamics due to two unique properties. Firstly, its creep strength is likely to be considerably smaller than that of (Mg,Fe)SiO₃ perovskite (Yamazaki and Karato 2001) and recent simulations suggest that perovskite will control the strength of the lower mantle while (Mg,Fe)O may accommodate most of the strain (Madi et al. 2005). Secondly, despite its cubic crystal structure, (Mg,Fe)O is likely to have large elastic anisotropy: according to first principles calculations, elastic anisotropy in (Mg,Fe)O is similar to that of orthorhombic (Mg,Fe)SiO₃ perovskite at lower mantle conditions (Karki et al. 1997; Wentzcovitch et al. 1998). Consequently, magnesiowüstite may contribute significantly to the development of seismic anisotropy (e.g. Karato 1998a, b).

Several studies have focused on texture development in periclase and magnesiowüstite. Merkel et al. (2002) deformed MgO at room temperature and high pressure. Yamazaki and Karato (2002), Stretton et al. (2001), and Heidelberg et al. (2003) deformed magnesiowüstite at moderate pressure and high temperature in torsion as well as axial compression. The results indicate that at low temperature $\{110\}\langle 1-10 \rangle$ is the only significantly active slip system, whereas at higher temperature $\{100\}$ and $\{111\}$ slips become equally active. In this report we investigate texture development in magnesiowüstite with compositions $(\text{Mg}_{0.4}\text{Fe}_{0.6})\text{O}$, $(\text{Mg}_{0.25}\text{Fe}_{0.75})\text{O}$, and

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(Mg_{0.1}Fe_{0.9})O in axial compression at high pressure and room temperature, measured in situ by synchrotron X-ray diffraction and a diamond anvil cell (DAC), with the goal to study the influence of iron content on texture development and elastic properties.

Experimental details

The starting materials were synthetic polycrystalline (Mg_{0.4}Fe_{0.6})O, (Mg_{0.25}Fe_{0.75})O, and (Mg_{0.1}Fe_{0.9})O that belong to the collection of magnesiowüstites synthesized and investigated by Bonczar and Graham (1982). They were prepared by reacting finely ground and well-mixed powders of oxides and iron in the appropriate proportions. To induce reaction, the palletized mixtures were sintered at 1,000°C for several hours in iron crucibles within an evacuated silica tube. Samples were then ground and the conversion to the B1 phase checked by X-ray diffraction. The chemical composition of the samples was determined using electron microprobe and wet chemical analyses. Before the present experiments, each sample was further ground in an corund mortar and loaded as a powder in the DAC. The grain size of the starting material, estimated in a stereomicroscope (Stemi SV11) with a zoom range 11:1 with a built in scale, was in the submicron range.

We performed X-ray diffraction experiments at the 16-IDB beamline of HPCAT at APS (Advanced Photon Source, Argonne National Laboratory) and at beamline 12.2.2, at ALS (Advanced Light Source, Lawrence Berkeley National Laboratory). For the experiments at HPCAT a monochromatic X-ray beam ($\lambda = 0.42275 \text{ \AA}$), approximately 15–17 μm in diameter, was focused on the sample either through a beryllium gasket for (Mg_{0.25}Fe_{0.75})O or amorphous boron and epoxy gasket for (Mg_{0.4}Fe_{0.6})O. Exposure time was 30 s for (Mg_{0.25}Fe_{0.75})O and 1 s for (Mg_{0.4}Fe_{0.6})O. (Mg_{0.1}Fe_{0.9})O samples were studied at beamline 12.2.2 at ALS using monochromatic X-ray beam ($\lambda = 0.4133 \text{ \AA}$ for the first run and $\lambda = 0.49593 \text{ \AA}$ for the second run). In this case, the 70 μm diameter beam was focused on the sample through a composite gasket made of an amorphous boron/epoxy disk confined within a kapton supporting ring (Merkel and Yagi 2005) with an exposure time of 240 s per image. Simple geometrical calculations indicate that for a beam of 15 and 70 μm diameter, the irradiated volume is on the order of 2.5×10^4 and $10^5 \mu\text{m}^3$, respectively. Therefore, each diffraction image includes a contribution from thousands of grains and is appropriate for texture refinement.

In all experiments X-ray diffraction was measured in radial geometry with the DAC compression axis perpendicular to the incident beam. Diffraction images were recorded with a MAR345 image-plate system (3,450×3,450 pixels). On those images, the intensity variations along Debye rings are indicative of preferred orientation while variations in diffraction angles are related to stress and elasticity (Fig. 1a, b). For the

measurements performed at the APS the detector to sample distance, calibrated using a Si standard, was 36.12 cm, thus optimizing the range of diffraction rings recorded for the opening cone of a standard symmetric piston-cylinder DAC to $2\theta_{\text{max}} = 18^\circ$. In the experiments performed at beamline 12.2.2 at ALS, the sample to detector distance, calibrated using a LaB₆ standard, was 29.69 cm, thus optimizing the range of diffractions recorded for the opening cone of a large-opening Mao-Bell type DAC to $2\theta_{\text{max}} = 26^\circ$. In all cases, hydrostatic pressures were estimated from the lattice parameter of a gold standard using the equation of state of gold (Anderson et al. 1989; Duffy et al. 1999). The uncertainty of the estimated pressures is $\pm 3 \text{ GPa}$ in average due to uncertainties in the equations of state of gold and nonhydrostatic stress conditions.

Data analysis

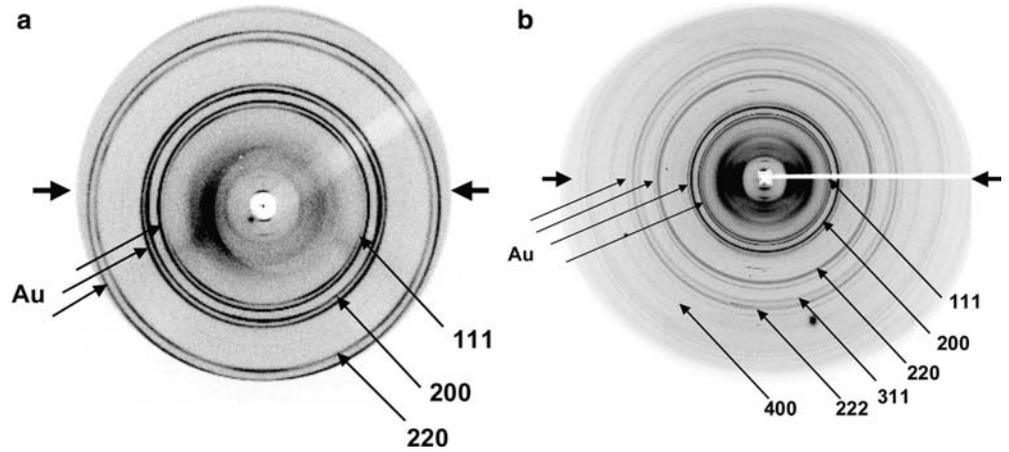
Rietveld analysis with MAUD

The X-ray diffraction images were analyzed with the Rietveld method using the software MAUD (Lutterotti et al. 1999; Lonardelli et al. 2005). This software refines instrumental, structural, microstructural parameters and texture from continuous diffraction spectra. Spectra were obtained by integrating the images over 5° azimuthal angular slices resulting in 72 spectra per image. In some of the experiments a few spectra had to be removed due to a saturation of the image because of a few large grains. Figure 2a, b shows representative spectra at different angles relative to the compression direction for (Mg_{0.4}Fe_{0.6})O and (Mg_{0.1}Fe_{0.9})O at 13 GPa.

For (Mg_{0.4}Fe_{0.6})O and (Mg_{0.25}Fe_{0.75})O we applied the Rietveld refinement method to all the azimuthal diffraction profiles with d -spacings ranging from 1.4 to 2.78 \AA which included 111, 200, and 220 (Fig. 2a). Diffraction profiles of (Mg_{0.1}Fe_{0.9})O were refined over a d -spacing range from 0.92 to 2.78 \AA (which includes six diffraction lines, 400, 222, 311, 220, 200, 111) (Fig. 2b). Instrument parameters (detector distance, beam center, peak profile, and detector tilt) were refined on the calibrant patterns. For each image, we then refined four background parameters per orientation, crystallographic parameters, stress, and texture with a composite sample of (Mg,Fe)O and gold. Stress was assumed to be triaxial with $2\sigma_{11} = 2\sigma_{22} = -\sigma_{33}$ (axial symmetry) and refined using a simple isotropic model based on Young's modulus and Poisson ratio.

In absence of experimental or theoretical data regarding the single-crystal elasticity of magnesiowüstite at high pressure we estimated the aggregate bulk and shear moduli of our sample using third order Eulerian strain equations (Davies 1974) based on the ambient pressure bulk and shear moduli of Jacobsen et al. (2002) and the pressure derivatives of Bonczar and Graham (1982). Once a satisfying fit between the calculated and measured d -spacings was achieved, we applied the tex-

Fig. 1 Representative diffraction patterns of magnesiowüstite, $(\text{Mg}_{0.4}\text{Fe}_{0.6})\text{O}$ (a) and $(\text{Mg}_{0.1}\text{Fe}_{0.9})\text{O}$ (b), both recorded at 13 GPa



ture model EWIMV that relies on the discrete tomographic method WIMV (Matthies and Vinel 1982) to deduce the orientation distribution function (ODF) of the sample. The main difference between EWIMV and WIMV is that in EWIMV the ODF cell path for each data point is computed explicitly for the true measurement angles, no longer requiring a regular grid coverage and interpolation. In this process, we chose a resolution of 10° and a projection tube radius of 20° . First, no sample symmetry was imposed. Under this approximation, we were able to obtain qualitatively reasonable pole figures from a single image (e.g. Ischia et al. 2005) indicating overall axial symmetry with some deviations that can be attributed to the minimal pole figure coverage. Figure 3a presents the 100 pole figure obtained in such refinement for $(\text{Mg}_{0.4}\text{Fe}_{0.6})\text{O}$ at 13 GPa. In a second cycle of refinements, cylindrical symmetry was imposed and Fig. 3b presents the symmetrical 100 pole figure obtained for the same conditions as Fig. 3a. After confirming axial symmetry we assumed it for all results presented in this paper.

At the end of the Rietveld refinement in MAUD the ODFs were exported and further smoothed in BEARTEX (Wenk et al. 1998) with a 11° Gauss filter to reduce stochastic effects. From this smoothed ODF we calculated inverse pole figures that represent the probability of finding crystal directions parallel to the compression direction. Densities are expressed in multiples of a random distribution (m.r.d.). For axially symmetric texture, inverse pole figures provide a complete texture description. The texture is also represented by some numerical parameters such as minimum and maximum of the ODF and the texture index $F2$ (Bunge 1982). The texture index is the integral over squared ODF values and thus it is a measure for the texture strength. Since all textures display a main concentration near 001 we also calculated how much grains with the compression direction parallel to 001 contribute to the overall texture. This was done by calculating the orientation volume of a sphere centered at Euler angles $\alpha=0^\circ$, $\beta=0^\circ$, $\gamma=0^\circ$ (Matthies convention) with a radius of 20° (program COMP in BEARTEX).

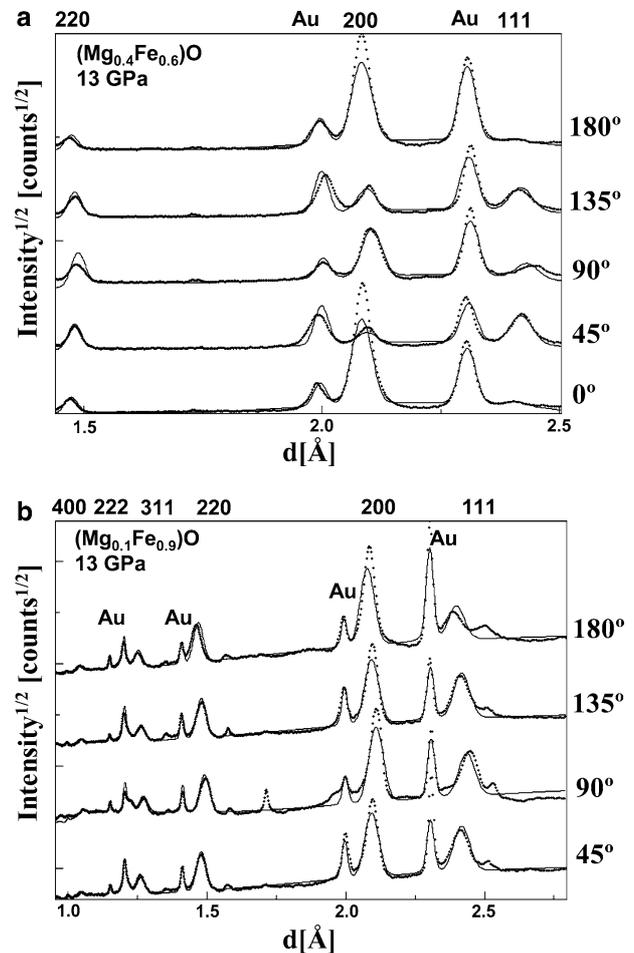
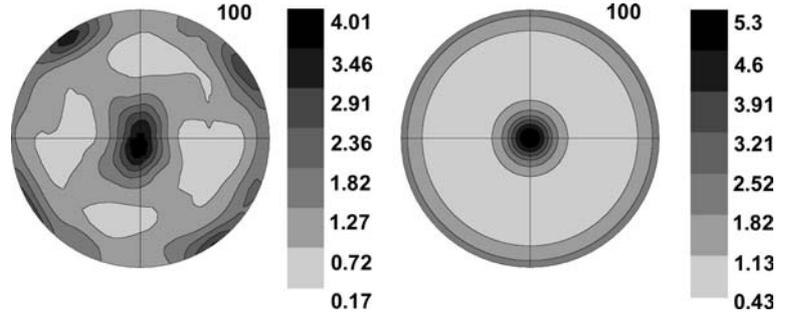


Fig. 2 Selected diffraction profiles for $(\text{Mg}_{0.4}\text{Fe}_{0.6})\text{O}$ (a) and $(\text{Mg}_{0.1}\text{Fe}_{0.9})\text{O}$ (b) at 13 GPa. *Dotted lines* are experimental data, *solid lines* are results from full profile refinement with the Rietveld technique. The profiles at 0° and 180° are in the compression direction and the profiles at 90° are in the extension direction. The 0° spectrum in **b** is not shown as it corresponds to the position of the beam-stop. Note the variations in peak positions and peak intensities with orientation, which are related to deviatoric stress and texture. A broadening of the peaks is observed. Diffraction lines of the sample are labeled in the figure. The diffraction peaks of the Au pressure standard are also indicated

Fig. 3 (100) pole figures of $(\text{Mg}_{0.4}\text{Fe}_{0.6})\text{O}$ smoothed with 11° Gauss filter without imposing symmetry **(a)** and with imposed cylindrical symmetry **(b)** at 13 GPa. Equal area projection. Four-fold symmetry in **a** is an artifact resulting from the pole figure coverage



Texture analysis with the single peak method

One of the major difficulties encountered in the refinement of the X-ray diffraction data of magnesio-wüstite is the presence of a large elastic strain anisotropy for different crystallographic directions. This reflects a strong directional dispersion of the Young's modulus $E(hkl)$. Based on data at ambient conditions (Jackson and Khanna 1990; Sinogeikin and Bass 2000; Jacobsen et al. 2002), the variation of Young's modulus between [100] and [111] ranges between 6 and 20% in the range of compositions investigated in this study. Because of this, texture analysis with the Rietveld method that uses an isotropic stress model was limited and we also applied a different approach based on single peak fit to extract texture and stress information.

In this second approach, we analyzed the variations in the position of the diffraction peaks and their intensity with the azimuthal angle δ by integrating intensities in 5° slices with the software FIT2D (Hammersley 1998). For each image, this produces 72 segments for δ between 0° and 360° . The patterns were then fitted individually assuming Gaussian peak profiles and a linear local background. The pole distance χ , which corresponds to the angle between the diffracting plane normal and the load axis, was calculated from

$$\cos \chi = \cos \theta \cos \delta, \quad (1)$$

where θ is the diffraction angle [see Fig. 3 in Merkel et al. (2002)].

Using this alternative approach each peak was fitted individually, and the diffraction intensities for each peak at all orientations could be satisfactorily reproduced. The extracted variations of diffraction intensity with orientation for (200), (220), and (111) were then used to calculate the ODF of the sample, assuming axial symmetry, with the WIMV algorithm of the BEARTEX package. Figure 4a–d presents a comparison for reflection (200) between the experimental intensities (circles and dots), those recalculated from the ODF fitted in the Rietveld package MAUD (dashed lines), and those recalculated from the ODF fitted in BEARTEX (solid lines). Overall, only small differences are observed.

Lattice strain analysis

The X-ray diffraction data were then analyzed to determine strain and stress using the procedure outlined in Singh et al. (1998) and briefly summarized here. Following the convention used in elasticity theory of single crystals (e.g. Nye 1960), the stress state at the center of the specimen is given by

$$\sigma_{ij} = \begin{pmatrix} \sigma_1 & 0 & 0 \\ 0 & \sigma_1 & 0 \\ 0 & 0 & \sigma_3 \end{pmatrix} = \begin{pmatrix} \sigma_P & 0 & 0 \\ 0 & \sigma_P & 0 \\ 0 & 0 & \sigma_P \end{pmatrix} + \begin{pmatrix} -t/3 & 0 & 0 \\ 0 & -t/3 & 0 \\ 0 & 0 & 2t/3 \end{pmatrix}, \quad (2)$$

where $(\sigma_3 - \sigma_1) \equiv t$. t is a measure of the deviatoric stress and is defined as uniaxial stress component (USC) following the nomenclature of Singh and Kennedy (1974) and Singh and Balasingh (1977). σ_P is the equivalent hydrostatic pressure.

The d -spacings for a given set of lattice planes measured by X-ray diffraction vary as a function of the angle χ between the compression direction and the diffracting plane normal according to the relation:

$$d_m(hkl) = d_p(hkl)[1 + (1 - 3 \cos^2 \chi)Q(hkl)], \quad (3)$$

where $d_m(hkl)$ is the measured d -spacing and $d_p(hkl)$ the d -spacing under the equivalent hydrostatic pressure σ_P and

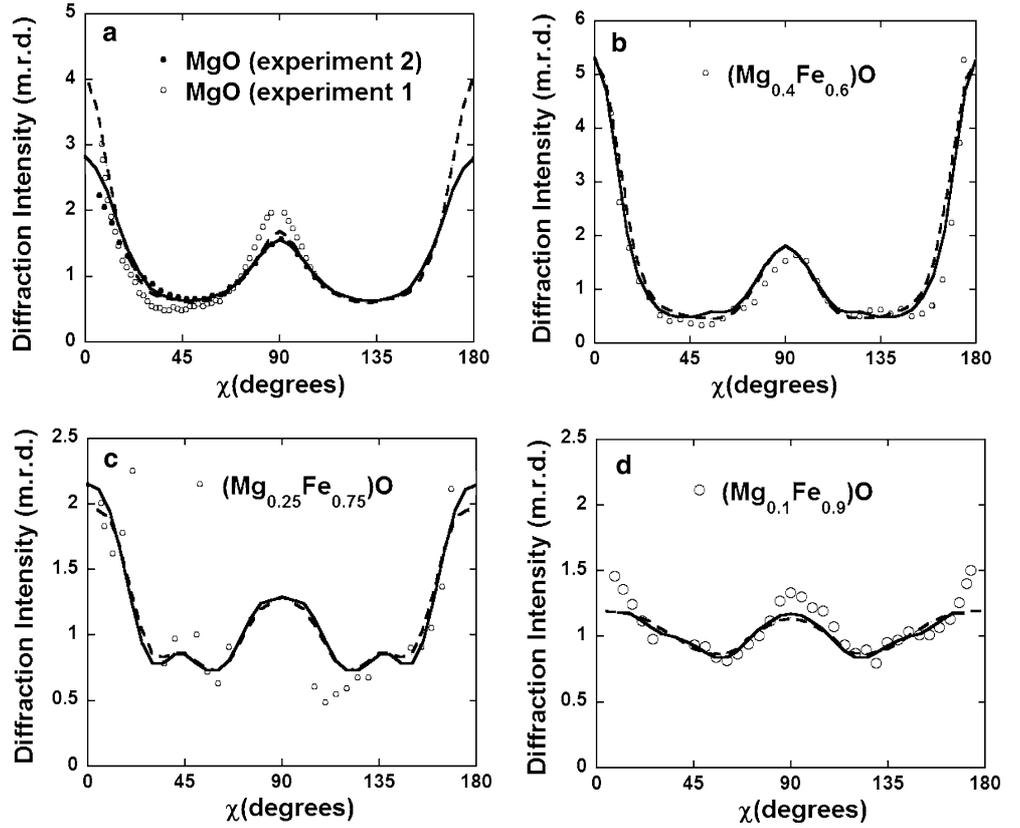
$$Q(hkl) = \frac{t}{3} \left[\frac{\alpha}{2G_R(hkl)} + \frac{1 - \alpha}{2G_V} \right], \quad (4)$$

where $G_R(hkl)$ and G_V are bounds to the shear modulus of the aggregate under the Reuss (iso-stress) and Voigt (iso-strain) approximations, respectively (Singh et al. 1998). The factor α which lies between 0 and 1 determines the relative weight of isostress (Reuss) and isostrain (Voigt) conditions. It specifies the degree of stress and strain continuity across grains in the sample.

For a cubic material, we have

$$(2G_V)^{-1} = \frac{5}{2} \frac{(S_{11} - S_{22})S_{44}}{[3(S_{11} - S_{12}) + S_{44}]} \quad (5)$$

Fig. 4 Variation of the diffraction intensities with χ for a pure MgO, **b** $(\text{Mg}_{0.4}\text{Fe}_{0.6})\text{O}$, **c** $(\text{Mg}_{0.25}\text{Fe}_{0.75})\text{O}$, and **d** $(\text{Mg}_{0.1}\text{Fe}_{0.9})\text{O}$ at about 13 GPa. *Open* (1. experiment) and *full circles* (2. experiment) are experimental data and *solid and dashed lines* are recalculated from the ODF fitted to the experimental intensities with WIMV as implemented in BEARTEX (*solid lines*) and ODF fitted with MAUD using the EWIMV algorithm (*dashed lines*), respectively



and

$$[2G_{\text{R}}(hkl)]^{-1} = S_{11} - S_{12} - 3S\Gamma(hkl), \quad (6)$$

where $S = (S_{11} - S_{12} - S_{44})/2$ is a measure of elastic anisotropy, and where the S_{ij} are the single crystal elastic compliances. The orientation factor $\Gamma(hkl) = \left(\frac{h^2k^2 + k^2l^2 + l^2h^2}{(h^2 + k^2 + l^2)^2} \right)$ ranges from 0 for the 200 line to a maximum of 1/3 for 111.

The axial stress, which represents a proxy for the material's strength (e.g. Singh et al. 1998; Kavner 2003) can be estimated from experimental results using the relation:

$$t = \sigma_3 - \sigma_1 \approx 6G \langle Q(hkl) \rangle, \quad (7)$$

where G is Hill's average of the Voigt and Reuss bounds to the shear modulus of the aggregate, and $\langle Q(hkl) \rangle$ is the average of the $\langle Q(hkl) \rangle$ deduced from the experimental data for the different diffraction lines (Singh et al. 1998).

The variations of the experimental d -spacings of the (111), (200), and (220) planes with χ upon compression and decompression were least-square fitted to Eq. 3 (Fig. 5a-c) in order to deduce $d_{\text{p}}(hkl)$ and $Q(hkl)$. In each case we allowed the maximum stress direction to be slightly shifted from $\delta = 0^\circ$ by introducing an offset in Eq. 3 as it could be observed in the experimental data that the d -spacings minima and maxima were not always perfectly aligned at 0° and 90° . This offset (on the order of $5-10^\circ$) could only be observed for lattice strains

measurements and is not observed in the sample texture (i.e. diffraction intensities). $Q(hkl)$ values were then plotted as a function of Γ and fitted to Eq. 4, which is linear in Γ for the cubic case (Fig. 6). $\langle Q(hkl) \rangle$, which represents an estimation of the strength of the sample normalized to the shear modulus ($\langle Q(hkl) \rangle = t/6G$, from Eq. 7) was calculated from the linear regression. Finally, the fitted slope (m_1) and the ordinate axis value (m_0) was used to calculate elastic anisotropy (A) (Fig. 6).

$$A = \left(\frac{2C_{44}}{C_{11} - C_{12}} \right) = \frac{3m_0}{3m_0 + m_1} \quad (8)$$

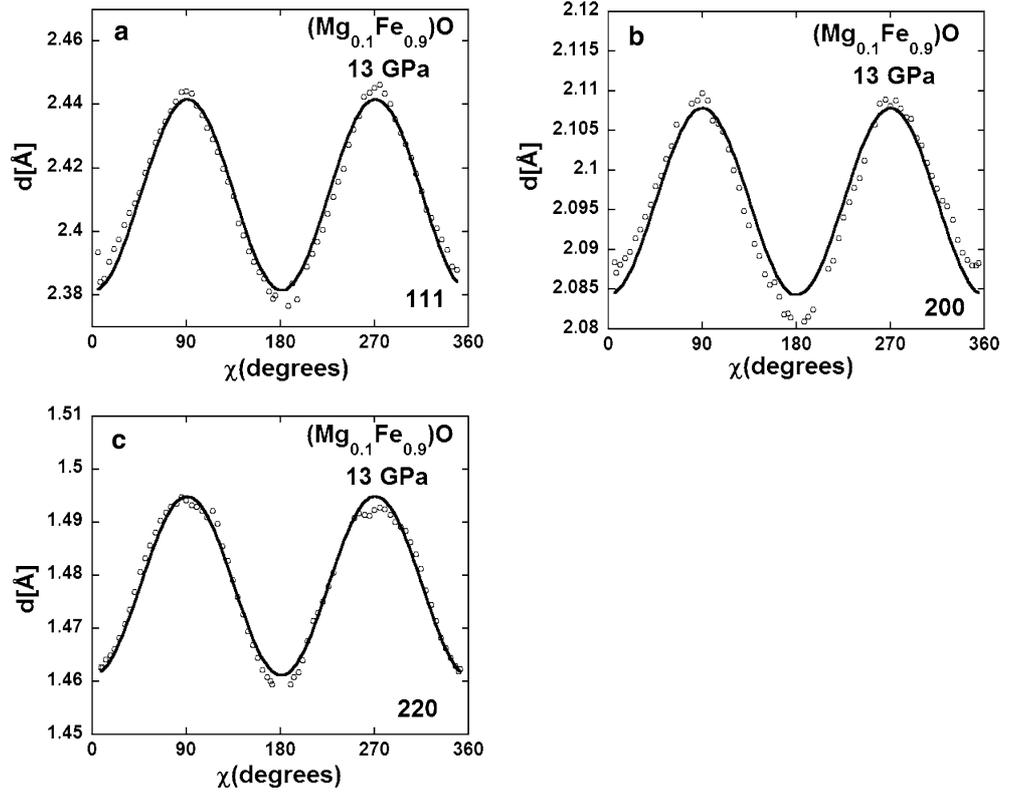
where the C_{ij} are the single crystal elastic moduli of the material.

Results

Rietveld analysis with MAUD

Lattice parameters of $(\text{Mg}_{0.40}\text{Fe}_{0.60})\text{O}$, $(\text{Mg}_{0.25}\text{Fe}_{0.75})\text{O}$, and $(\text{Mg}_{0.10}\text{Fe}_{0.90})\text{O}$ were refined for each pressure and composition (Table 1). We have fixed the uncertainty on the refined unit cell parameters to an average value of $\pm 0.005 \text{ \AA}$, consistent with the precision of peak position determination in our X-ray diffraction experiments [$\sigma(d_{\text{hkl}}) \sim 10^{-4}$ at beamline 12.2.2 of the Advanced Light Source]. The estimated standard deviations on the unit cell parameter calculated in the Rietveld refinement appear largely underestimated because of the many

Fig. 5 $d_m(111)$ (a), $d_m(200)$ (b), and $d_m(220)$ (c) versus χ for $(\text{Mg}_{0.1}\text{Fe}_{0.9})\text{O}$ at 13 GPa. Circles are experimental data and solid lines are fits of lattice strain equations



assumptions involved in this whole spectrum fit procedure. The pressure dependence (V/V_0) that we determined for $(\text{Mg}_{0.40}\text{Fe}_{0.60})\text{O}$, $(\text{Mg}_{0.25}\text{Fe}_{0.75})\text{O}$, and $(\text{Mg}_{0.10}\text{Fe}_{0.90})\text{O}$ is in agreement, within estimated uncertainties, with the existing 300 K isothermal equation of state for similar compositions (Jacobsen et al. 2002; Mao et al. 2002) (Fig. 7a–c).

The principal components of the deviatoric stress tensor (σ_{11} , σ_{22} , σ_{33}) were evaluated using a simple tri-axial stress model as part of MAUD refinement procedure, based on estimated bulk and shear moduli (see “Data analysis”). To account for the typical geometry of stress in the DAC, σ_{33} was refined with $\sigma_{22} = \sigma_{11}$ and $\sigma_{11} + \sigma_{22} + \sigma_{33} = 0$. The maximum principal component

Table 1 Unit cell parameters of $(\text{Mg}_{0.4}\text{Fe}_{0.6})\text{O}$, $(\text{Mg}_{0.25}\text{Fe}_{0.75})\text{O}$, $(\text{Mg}_{0.1}\text{Fe}_{0.9})\text{O}$ (run 1) $(\text{Mg}_{0.4}\text{Fe}_{0.6})\text{O}$ (run 2). Principal stresses components obtained from the MAUD refinements

| | a_{MW} [Å] (± 0.005) | V/V_0 | σ_{33} (GPa) | Gold a (Å) | P (GPa) |
|---|-------------------------------------|---------|---------------------|--------------|-----------|
| Sample $(\text{Mg}_{0.4}\text{Fe}_{0.6})\text{O}$ | | | | | |
| 4060_2_004 | 4.243 | 0.980 | −0.5 | 4.048 | 4 |
| 4060_2_005 | 4.215 | 0.961 | −1.5 | 4.018 | 9 |
| 4060_2_006 | 4.197 | 0.948 | −1.7 | 4.002 | 11 |
| 4060_2_009 | 4.175 | 0.934 | −1.3 | 3.974 | 16 |
| 4060_2_011 | 4.144 | 0.913 | −2.5 | 3.948 | 22 |
| 4060_2_013 | 4.116 | 0.895 | −2.9 | 3.925 | 27 |
| 4060_2_015 | 4.099 | 0.883 | −2.6 | 3.912 | 30 |
| 4060_2_016 | 4.086 | 0.875 | −2.6 | 3.901 | 33 |
| 4060_2_018 | 4.075 | 0.868 | −3.0 | 3.884 | 37 |
| Sample $(\text{Mg}_{0.25}\text{Fe}_{0.75})\text{O}$ | | | | | |
| 2575_003 | 4.265 | 0.988 | −0.1 | 4.067 | 2 |
| 2575_017 | 4.243 | 0.973 | −0.7 | 4.048 | 4 |
| 2575_019 | 4.230 | 0.964 | −1 | 4.038 | 6 |
| 2575_027 | 4.188 | 0.935 | −1.3 | 4.011 | 10 |
| 2575_028 | 4.162 | 0.918 | −1.4 | 3.978 | 16 |
| Sample $(\text{Mg}_{0.1}\text{Fe}_{0.9})\text{O}$ | | | | | |
| 1090_005 | 4.258 | 0.974 | −1 | 4.047 | 4 |
| 1090_006 | 4.197 | 0.933 | −1.4 | 3.994 | 13 |
| 1090_007 | 4.167 | 0.913 | −1.8 | 3.968 | 18 |
| Sample $(\text{Mg}_{0.1}\text{Fe}_{0.9})\text{O}$ | | | | | |
| 1090_2_006 | 4.262 | 0.977 | −1.1 | | 4 |
| 1090_2_008 | 4.191 | 0.929 | −1.6 | | 12 |

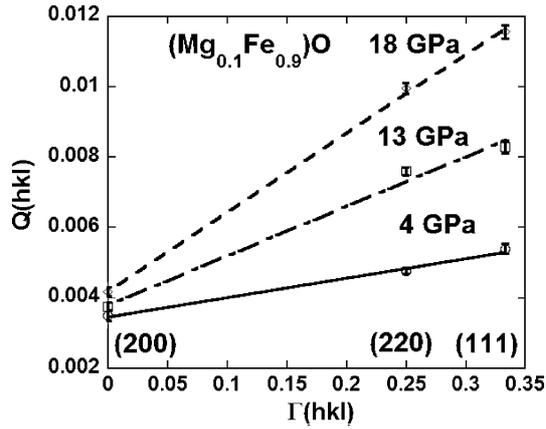


Fig. 6 $Q(hkl)$ versus $\Gamma(hkl)$ for $(Mg_{0.1}Fe_{0.9})O$ at 4, 13, and 18 GPa

of the deviatoric stress tensor (σ_{33}) obtained from the refinements tends to increase and later saturate with increasing pressure and range from 0 to 3 GPa (Tables 1).

Our texture analysis with the Rietveld approach shows that $(Mg_{0.4}Fe_{0.6})O$, $(Mg_{0.25}Fe_{0.75})O$, and $(Mg_{0.1}Fe_{0.9})O$ deformed in compression in the DAC mainly have a [100] texture as displayed in inverse pole figures (Fig. 8). Table 2 presents the minima and maxima of all ODFs calculated in this study. The texture strength $F2$ ranges between 1.7 and 2.3 for $(Mg_{0.4}Fe_{0.6})O$, 1.1 and 1.3 for $(Mg_{0.25}Fe_{0.75})O$, and 1.0 and 1.02 for $(Mg_{0.1}Fe_{0.9})O$. The texture index of MgO

determined by Merkel et al. (2002) in a comparable pressure range (0–35 GPa) ranges between 1 and 2.7 in two different experiments.

The volume fraction of the main texture component at [100] (as defined in “Experimental details”) for all the samples and all pressures is reported in Table 2 and plotted as a function of pressure in Fig. 9. Its pressure and compositional dependencies are in good agreement with those observed for both $F2$ index and ODF maximum values. The combination of these results shows that the texture strength of magnesiowüstites (a) increases and later saturates with increasing pressure, and (b) weakens with increasing Fe content. One of the experiments (experiment 2) performed on MgO by Merkel et al. (2002) shows a weaker pressure dependence of texture strength. One possible reason of this behavior could be found in the sample arrangement used in this second experiment (MgO was loaded as thermal insulator and pressure transmitting medium around Fe).

Texture analysis performed independently by individual fit of peaks confirms that the main texture of magnesiowüstites is at [100]. Figure 4a–d presents the intensities of diffraction peak 200 of MgO, $(Mg_{0.4}Fe_{0.6})O$, $(Mg_{0.25}Fe_{0.75})O$, and $(Mg_{0.1}Fe_{0.9})O$ as a function of the angle χ at comparable pressures. The overall variation of amplitude of the diffraction intensities with χ decreases with increasing iron content, such that $(Mg_{0.1}Fe_{0.9})O$ shows a nearly random distribution. The inverse pole figures obtained with the whole spectrum fit approach (Rietveld method) (Fig. 8a.1, b.1)

Fig. 7 V/V_0 versus pressure for $(Mg_{0.4}Fe_{0.6})O$ (a), $(Mg_{0.25}Fe_{0.75})O$ (b) and $(Mg_{0.1}Fe_{0.9})O$ (c)

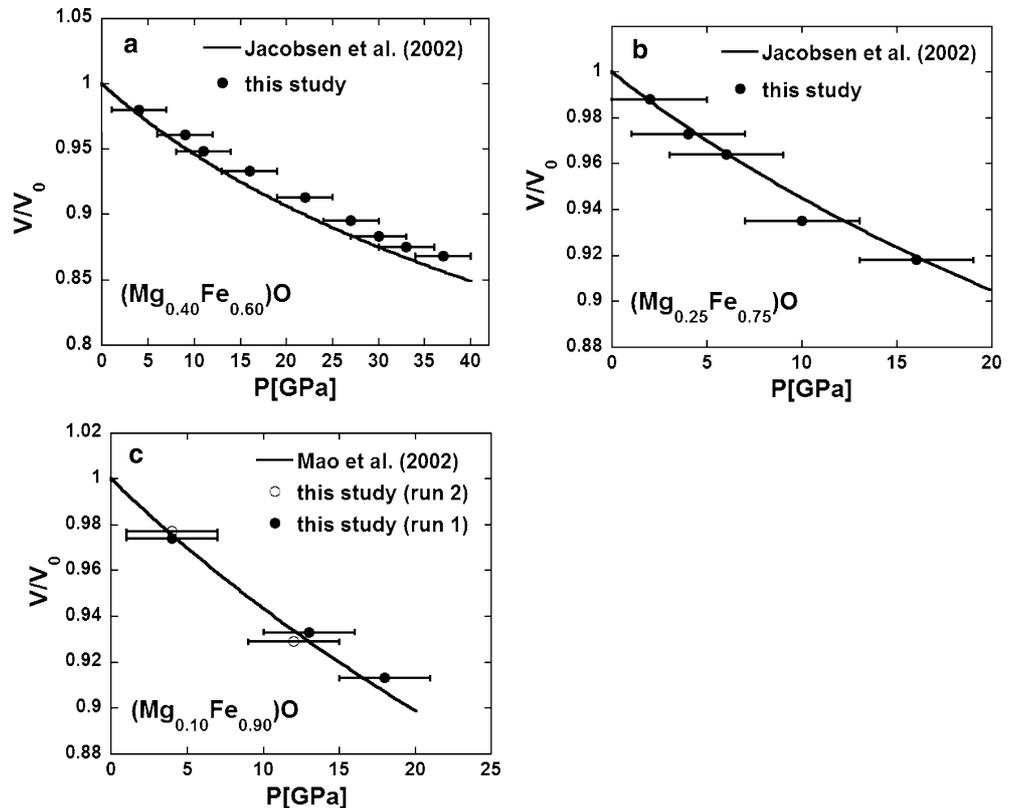


Table 2 Texture index F_2 , minima and maxima of the ODF, and principal pole figures 100 with imposing cylindrical symmetry. ($\text{Mg}_{0.4}\text{Fe}_{0.6}\text{O}$), ($\text{Mg}_{0.25}\text{Fe}_{0.75}\text{O}$), ($\text{Mg}_{0.1}\text{Fe}_{0.9}\text{O}$) (run 1), and ($\text{Mg}_{0.1}\text{Fe}_{0.9}\text{O}$) (run 2)

| | ODF ^a (m.r.d.) | | ODF ^b (m.r.d.) | | F_2 | Vol% component |
|---|---------------------------|------|---------------------------|------|-------|----------------|
| | Min | Max | Min | Max | | |
| Sample ($\text{Mg}_{0.4}\text{Fe}_{0.6}\text{O}$) | | | | | | |
| 4060_2_004 | 0.35 | 4.25 | | | | 7.76 |
| 4060_2_005 | 0.41 | 4.32 | 0.39 | 3.96 | 1.44 | 11.18 |
| 4060_2_006 | 0.32 | 3.08 | 0.79 | 1.23 | 1.47 | 11.65 |
| 4060_2_009 | 0.37 | 5.23 | 0.30 | 6.03 | 1.96 | 15.30 |
| 4060_2_011 | 0.04 | 7.82 | 0.41 | 5.21 | 1.91 | 14.93 |
| 4060_2_013 | 0.26 | 5.82 | 0.38 | 5.05 | 1.89 | 15.05 |
| 4060_2_015 | 0.23 | 5.73 | 0.43 | 5.44 | 1.89 | 14.36 |
| 4060_2_016 | 0.31 | 5.27 | 0.35 | 6.16 | 2.00 | 15.92 |
| 4060_2_018 | 0.20 | 6.89 | 0.44 | 5.87 | 2.07 | 16.40 |
| Sample ($\text{Mg}_{0.25}\text{Fe}_{0.75}\text{O}$) | | | | | | |
| 2575_003 | 0.53 | 2.06 | 0.47 | 1.73 | 1.04 | 5.19 |
| 2575_017 | 0.36 | 1.99 | 0.43 | 2.25 | 1.16 | 8.18 |
| 2575_019 | 0.33 | 2.30 | 0.38 | 2.24 | 1.15 | 8.32 |
| 2575_027 | 0.36 | 2.15 | 0.40 | 2.17 | 1.18 | 7.70 |
| 2575_028 | 0.62 | 1.79 | 0.33 | 1.60 | 1.13 | 5.99 |
| Sample ($\text{Mg}_{0.1}\text{Fe}_{0.9}\text{O}$) | | | | | | |
| 1090_005 | 0.83 | 1.13 | 0.90 | 1.13 | 1.00 | 1.42 |
| 1090_006 | 0.82 | 1.14 | 0.79 | 1.23 | 1.01 | 2.12 |
| 1090_007 | 0.83 | 1.13 | 0.63 | 1.17 | 1.02 | 2.91 |
| Sample ($\text{Mg}_{0.1}\text{Fe}_{0.9}\text{O}$) | | | | | | |
| 1090_2_006 | 0.75 | 1.14 | 0.92 | 1.09 | 1.00 | 0.82 |
| 1090_2_008 | 0.48 | 1.35 | 0.59 | 1.21 | 1.02 | 3.27 |

^aMinima and maxima of the ODF recalculated with WIMV as implemented in BEARTEX from the single peak fits

^bMinima and maxima of the ODF recalculated from MAUD using the EWIMV algorithms. The ODF were smoothed using a Gaussian filter of width of 1-1°. The radius of the texture components were set to 20°

confirm the results obtained in the inverse pole figures recalculated from the individual peak fits are shown in Fig. 8a.2, b.2.

Elastic anisotropy and stress

The axial stress component in the sample can be calculated from the variations of the d -spacings with orientation in the diffraction pattern. In the present analysis, we calculated the average axial stress component $\langle t \rangle$, defined as $6G\langle Q(hkl) \rangle$, where G is the average shear modulus of the aggregate under the Hill approximation. In order to assess effects of plastic deformation, we also evaluated individual stress components using Eq. 4 and the Reuss approximation for (200), (111), and (222) when data for (111) was not available.

The value of $\langle Q(hkl) \rangle$ increases with pressure from 0.03 to 0.06 in ($\text{Mg}_{0.40}\text{Fe}_{0.60}\text{O}$), 0.02–0.05 in ($\text{Mg}_{0.25}\text{Fe}_{0.75}\text{O}$), 0.03–0.06 in ($\text{Mg}_{0.10}\text{Fe}_{0.90}\text{O}$) (run 1), and 0.03–0.05 in ($\text{Mg}_{0.10}\text{Fe}_{0.90}\text{O}$) (run 2). In general, we find that stress ($6G\langle Q(hkl) \rangle$) increases and later saturates with increasing pressure. It is higher for the MgO samples measured by Merkel et al. (2002) than for the present magnesiowüstite samples (Fig. 10a; Table 3). For MgO, stresses measured on (222) tend to be larger than stresses on (200). For other compositions, stresses measured on (200) tend to be larger than those measured on (111).

The elastic anisotropy factor A for ($\text{Mg}_{0.40}\text{Fe}_{0.60}\text{O}$), ($\text{Mg}_{0.25}\text{Fe}_{0.75}\text{O}$), and ($\text{Mg}_{0.10}\text{Fe}_{0.90}\text{O}$) decreases with increasing pressure and increasing iron content (Fig. 10b; Table 3). A is the ratio of the shear moduli in the (100) and (110) planes in the [100] direction. If $A = 1$

the material is elastically isotropic. We find a change of sign of the anisotropy factor of magnesiowüstite at a pressure that depends on composition, leading to an interchange in the direction of fastest and slowest compression and shear wave propagation. For instance, at low pressure, compression waves travel faster along $\langle 111 \rangle$, at high pressure $\langle 100 \rangle$ is the fastest direction. The pressure of reversal obtained from the radial diffraction data is in reasonable agreement with simple models based on third order Eulerian strain equations (Davies 1974) using the ambient pressure elastic moduli of Jacobsen et al. (2002) and the pressure derivatives of Bonczar and Graham (1982) (Fig. 10b).

Discussion

Lattice preferred orientation

For all examined compositions, our experiments document the development of a main texture component at [100] with increasing pressure. The [100] texture is due to $\{110\}\langle 1-10 \rangle$ slip as explored previously (Wenk et al. 1989; Merkel et al. 2002). If $\{110\}\langle 1-10 \rangle$, $\{111\}\langle 1-10 \rangle$, and $\{100\}\langle 011 \rangle$ slip systems were all active simultaneously, the maximum in the inverse pole figure would be a girdle between [110] and [100]. Textures we obtain in ($\text{Mg}_{0.1}\text{Fe}_{0.9}\text{O}$) are particularly weak and sometimes show other maxima than [100]. However, the intensity of those secondary maxima is weak and depends on the fitting method.

In order to study the influence of iron on (Mg,Fe)O we compared our results with those of pure MgO as a reference under similar experimental conditions (Merkel

Table 3 Values of effective stresses deduced from the X-ray data on (111) and (200), mean stress value ($\langle t \rangle = 6G\langle Q(hkl) \rangle$), and elastic anisotropy for all compositions studied here along with the MgO results of Merkel et al. (2002). For MgO, minima and maxima of

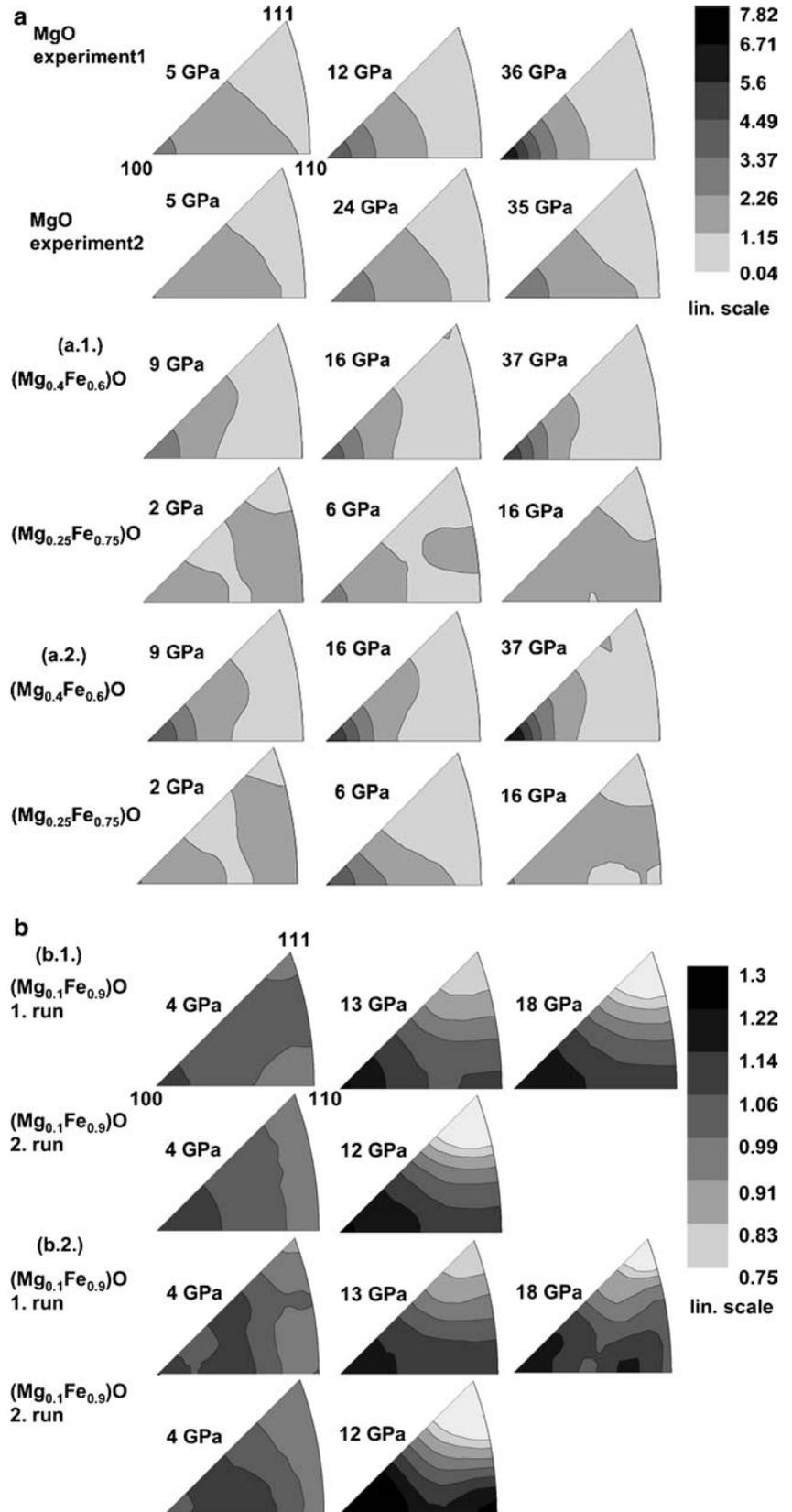
the ODF were deduced from experimental ODF calculated with WIMV and smoothed using a Gaussian filter of width of 11° , as implemented in BEARTEX

| P (GPa) | $t(222)$ | $t(200)$ | $\langle t \rangle$ | $A[2C_{44}/(C_{11}-C_{12})]$ | ODF | | $F2$ |
|---|----------|----------|---------------------|------------------------------|------|------|------|
| | | | | | Min | Max | |
| MgO (experiment 1) | | | | | | | |
| 1.9 | 1.5 | 1.7 | 1.7 | 1.674 | 0.91 | 1.10 | 1.00 |
| 2.4 | 2.0 | 3.0 | 2.7 | 1.681 | 0.89 | 1.19 | 1.01 |
| 3.1 | 3.0 | 2.7 | 3.1 | 1.508 | 0.87 | 1.26 | 1.01 |
| 5.3 | 5.6 | 5.1 | 5.5 | 1.252 | 0.70 | 2.20 | 1.11 |
| 7.0 | 5.9 | 5.5 | 6.0 | 1.058 | 0.63 | 2.96 | 1.24 |
| 9.5 | 7.1 | 6.1 | 6.8 | 0.984 | 0.52 | 3.78 | 1.41 |
| 11.6 | 7.0 | 6.0 | 6.6 | 0.987 | 0.52 | 3.81 | 1.41 |
| 20.3 | 8.1 | 6.4 | 7.1 | 0.885 | 0.22 | 6.84 | 2.17 |
| 30.3 | 10.3 | 9.6 | 9.5 | 0.912 | 0.27 | 6.46 | 2.25 |
| 35.5 | 9.4 | 9.5 | 8.6 | 1.071 | 0.27 | 6.46 | 2.25 |
| MgO (experiment 2) | | | | | | | |
| 1.5 | 2.0 | 2.0 | 2.1 | 1.371 | | | |
| 5.1 | 4.8 | 3.8 | 4.3 | 0.991 | 0.71 | 2.00 | 1.09 |
| 7.5 | 8.6 | 6.0 | 7.5 | 0.910 | 0.61 | 2.15 | 1.14 |
| 11.0 | 9.8 | 7.0 | 8.5 | 0.869 | 0.53 | 2.76 | 1.26 |
| 17.4 | 9.5 | 6.8 | 8.0 | 0.870 | 0.52 | 2.78 | 1.26 |
| 18.3 | 8.7 | 6.9 | 7.7 | 0.931 | 0.48 | 2.57 | 1.24 |
| 20.2 | 9.0 | 7.2 | 7.8 | 0.960 | 0.40 | 2.34 | 1.21 |
| 23.0 | 10.5 | 7.3 | 8.6 | 0.844 | 0.53 | 2.58 | 1.24 |
| 23.9 | 9.1 | 7.6 | 8.3 | 0.836 | 0.46 | 2.80 | 1.27 |
| 27.0 | 9.6 | 7.9 | 8.4 | 0.852 | 0.40 | 2.74 | 1.26 |
| 29.8 | 10.6 | 9.5 | 9.7 | 0.895 | 0.47 | 2.81 | 1.27 |
| P (GPa) | $t(111)$ | $t(200)$ | $\langle t \rangle$ | $A[2C_{44}/(C_{11}-C_{12})]$ | | | |
| (Mg _{0.4} Fe _{0.6})O | | | | | | | |
| 4 | 2.1 | 1.9 | 2.1 | 0.923 | | | |
| 9 | 1.9 | 1.9 | 2.0 | 0.810 | | | |
| 11 | 2.3 | 2.3 | 2.3 | 0.739 | | | |
| 16 | 1.8 | 1.9 | 1.9 | 0.637 | | | |
| 22 | 3.2 | 3.7 | 3.3 | 0.572 | | | |
| 27 | 3.7 | 4.7 | 3.9 | 0.530 | | | |
| 30 | 2.8 | 4.3 | 3.3 | 0.578 | | | |
| 33 | 4.5 | 5.6 | 4.7 | 0.412 | | | |
| 37 | 4.2 | 6.0 | 4.7 | 0.406 | | | |
| (Mg _{0.25} Fe _{0.75})O | | | | | | | |
| 4 | 0.9 | 1.1 | 1.0 | 1.152 | | | |
| 6 | 1.2 | 1.5 | 1.4 | 0.953 | | | |
| 10 | 2.4 | 2.7 | 2.5 | 0.676 | | | |
| 16 | 2.2 | 3.2 | 2.6 | 0.588 | | | |
| (Mg _{0.1} Fe _{0.9})O (run 1) | | | | | | | |
| 4 | 1.6 | 1.4 | 1.5 | 0.651 | | | |
| 13 | 1.9 | 2.1 | 2.0 | 0.449 | | | |
| 18 | 2.2 | 2.8 | 2.5 | 0.360 | | | |
| (Mg _{0.1} Fe _{0.9})O (run 2) | | | | | | | |
| 4 | 1.9 | 1.7 | 1.8 | 0.658 | | | |
| 12 | 1.3 | 1.9 | 1.5 | 0.410 | | | |

et al. 2002). Indeed, we find that pure MgO has the same main texture component at [100] as our magnesiowüstite samples (Fig. 8). The texture strength tends to decrease

with increasing iron content (Table 2; Figs. 8, 9). With increasing iron content, bonding becomes less ionic and the strength decreases. This could result in lower

Fig. 8 Inverse pole figures for different compositions in the MgO–FeO series: **a.1** MgO (experiment 1 and 2), $(\text{Mg}_{0.4}\text{Fe}_{0.6})\text{O}$, $(\text{Mg}_{0.25}\text{Fe}_{0.75})\text{O}$ recalculated from MAUD using EWIMV algorithms. **a.2** The same as **a.1** recalculated with WIMV as implemented in BEARTEX from the single peak fits. **b.1** Inverse pole figures for $(\text{Mg}_{0.1}\text{Fe}_{0.9})\text{O}$ (runs 1 and 2) from MAUD using the EWIMV algorithms. **b.2** The same as **b.1** recalculated with WIMV as implemented in BEARTEX from the single peak fits. Because of the weaker texture, the $(\text{Mg}_{0.1}\text{Fe}_{0.9})\text{O}$ composition is presented with a different gray scale. Equal projection area



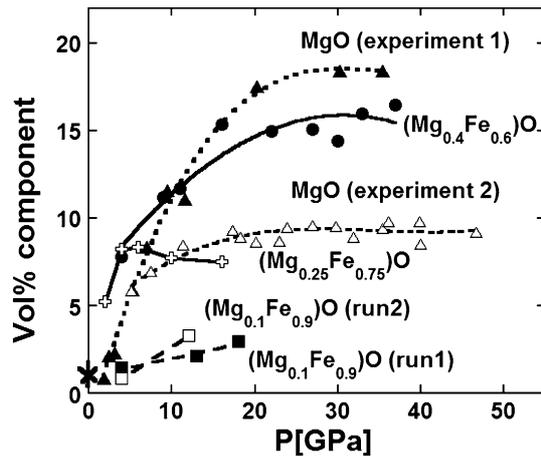


Fig. 9 100 texture component (vol%) versus pressure for different compositions in the MgO–FeO solid solution series. The radius of the spherical texture component was chosen at 20°. Data for $(\text{Mg}_{0.1}\text{Fe}_{0.9})\text{O}$ (open and closed squares), $(\text{Mg}_{0.25}\text{Fe}_{0.75})\text{O}$ (crosses), $(\text{Mg}_{0.4}\text{Fe}_{0.6})\text{O}$ (circles), and MgO (open and closed triangles) are shown in the figure. Lines are guide to the eye through the experimental data and the star symbol indicates the value that would be obtained for a random texture

stacking fault energy for the slip system $\{110\}\langle 1\bar{1}0\rangle$ (Miranda and Scandolo 2005). Another explanation could be that in iron-rich, more compressible, magnesiowüstites a larger portion of strain is accommodated elastically. The observed broadening of the diffraction lines in all three magnesiowüstite compositions with increasing pressure, may also be an expression of such heterogeneous elastic strain, probably because of higher dislocations densities.

$(\text{Mg}_{0.80}\text{Fe}_{0.20})\text{O}$ up to 300 MPa and 1,400 K by Stretton et al. (2001) show [100] and [110] texture in analyses using the EBSD technique. The samples at higher temperature (1,400 K) and larger grain size and those at lower temperature (1,200 K) and smaller grain size, show a stronger texture at [110], which broadens towards [100] with a minimum at [111]. The texture is only slightly stronger for the sample deformed at lower temperature (1,200 K). This suggests that temperature plays a relevant role on deformation mechanisms of magnesiowüstite, similar to isostructural halite (Carter and Heard 1970).

Heidelbach et al. (2003) studied texture evolution as a function of strain, observing a change from a dislocation creep texture at low strains to a recrystallization texture at high strains. With increasing strain the texture components in the following order become more evident: $\{111\}\langle 110\rangle$, $\{100\}\langle 110\rangle$ and finally at the highest strains the texture component $\{112\}\langle 110\rangle$. The overall observed texture strengths in this study are relatively low compared to the high strains. This is explained with the subgrain rotation recrystallization, weakening strong maxima due to rotation of the crystals.

Another aspect that has to be considered in order to understand the deformation behavior of magnesiowüstites is the complex defect structure associated

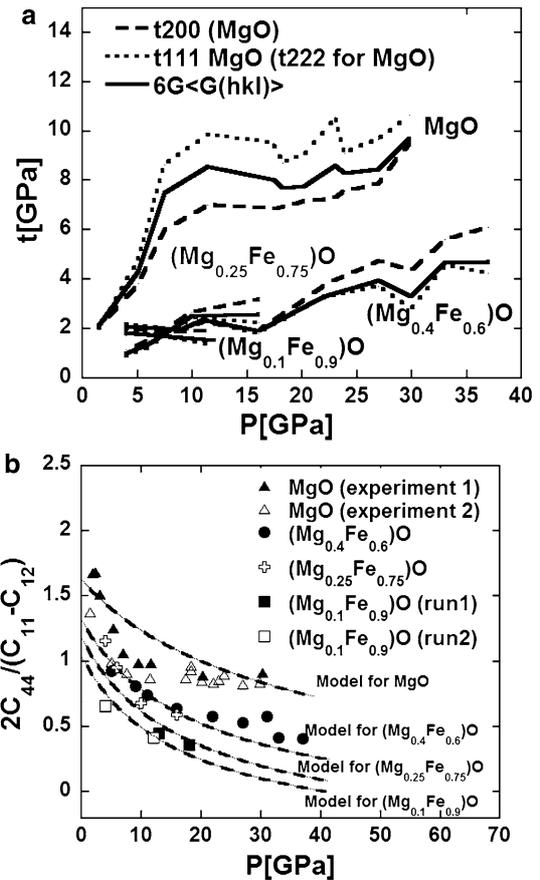


Fig. 10 **a** Mean stress ($6G\langle Q(hkl)\rangle$), and effective stresses deduced from measurements on (111) and (200) versus pressure for all samples as a function of pressure. For the measurements on MgO, effective stresses were calculated on (222) and (200) as no data for (111) was available. For clarity, measurements of MgO (experiment 1) and $(\text{Mg}_{0.1}\text{Fe}_{0.9})\text{O}$ (run 1) were not included in the figure but they do not differ significantly. **b** Elastic anisotropy in magnesiowüstite expressed with the Zener ratio $A = 2C_{44}/(C_{11} - C_{12})$ as a function of pressure. Results deduced from the X-ray data under the Reuss limit are compared with models based on data from literature (Jackson and Khanna 1990; Sinogeikin and Bass 2000; Jacobsen et al. 2002, 2004) (dashed lines)

with very high Fe contents. The less ionic bonding strength in $(\text{Mg,Fe})\text{O}$ and the probably existing vacancies in the $(\text{Mg,Fe})\text{O}$ crystal structure (Jacobsen et al. 2002) might increase the density of dislocations (Luecke and Kohlstedt 1988; Pulliam 1963; Heidelbach et al. 2003).

Finally, we cannot neglect that grain size, grain shape, grain distribution, configuration of the existing phases, deformation history, etc. in the sample can play a role in the deformation process.

Orientation of stress in the DAC

For each pressure, we could observe a slight offset in the maximum stress direction in the measured d -spacings. This indicates that the maximum stress applied to the sample is not always aligned with the anvil directions but

can be shifted by a few degrees. This is probably related to details of the experimental setup such as diamond alignment and sample loading. This has been also observed on other samples in DAC radial diffraction experiments (Merkel et al. 2005; S. Merkel and T. Yagi, submitted for publication). Interestingly, this offset is not observed in the measured diffraction intensities, indicating that the texture and plastic deformation of the sample is indeed axisymmetric. It should be noted that this can be problematic for whole profile Rietveld refinements in MAUD as this stress geometry cannot be readily refined.

Effect of plasticity on the deduced anisotropy

In order to better quantify the combined effects of pressure increase and Fe–Mg substitution on the elastic anisotropy in magnesiowüstites we have compared our experimental values of the anisotropy factor (A), with model anisotropies calculated from low-pressure single-crystal elasticity data (Jackson and Khanna 1990; Sinogeikin and Bass 2000; Jacobsen et al. 2002, 2004) extrapolated using Eulerian finite strain equations (Davies 1974). The model agrees reasonably well with our experimental results (Fig. 10b). These results indicate that the lattice strain theory (Singh et al. 1998) can be applied to obtain a first estimate of the anisotropies of cubic materials.

In recent years, models based on a combination of polycrystal plasticity and lattice strain theories predicted that for materials such as MgO, stresses deduced from lattice strain measurements on (111) should become over 50% greater than deduced from (200), once plastic deformation is activated (Weidner et al. 2004). This model appears to agree with the results on MgO of Merkel et al. (2002). However, it contradicts our measurements for iron-rich magnesiowüstites for which the stress measured on (200) is larger than that measured on (111) (Fig. 10a). Our experiments do indicate an (hkl)-dependence of stress but much lower than predicted by Weidner et al. (2004). It should be noted that elastoplastic models such as those used by Weidner et al. (2004) and Li et al. (2004) depend on a number of parameters, such as the combination of active slip system and the elastic anisotropy of the material. Our results indicate that the combination of parameters used in those models was optimized to match the data measured for MgO. In particular, elastic anisotropy in magnesiowüstite changes greatly with increasing iron content and therefore the models of Weidner et al. (2004) and Li et al. (2004) may not apply directly. The evolution of elastic anisotropy with pressure and composition deduced from this study is in fairly good agreement with other techniques (Fig. 10b). This indicates that the radial diffraction data can be used to extract trends of variation of anisotropies with composition. However, as pointed out by Weidner et al. (2004) and Li et al. (2004), the absolute values obtained should be treated with caution.

Conclusion

Axial deformation experiments on $(\text{Mg}_{0.4}\text{Fe}_{0.6})\text{O}$, $(\text{Mg}_{0.25}\text{Fe}_{0.75})\text{O}$, and $(\text{Mg}_{0.1}\text{Fe}_{0.9})\text{O}$ with the DAC show the development of a [100] texture. The comparison between the lattice preferred orientation in our (Mg,Fe)O samples with those observed in pure MgO, both at room temperature, suggests that the $\{1-10\}\langle 110\rangle$ is the main active slip system at these conditions. However, the weakening of texture with increasing iron content is significant.

Stress measured in our magnesiowüstite samples is significantly lower than previously measured in MgO (Merkel et al. 2002). Contrary to models based on polycrystal plasticity and lattice strain theories and applied to MgO (Weidner et al. 2004), we find that for iron-rich compositions, the effective stress measured on (111) is lower than that measured on (200). Elastoplastic models such as the ones used in Weidner et al. (2004) and Li et al. (2004) depend on a number of parameters, such as the combination of active slip system and the elastic anisotropy of the material, and our results underline that they can only be used to interpret and analyze the experiments they were designed for.

Finally, the elastic anisotropy deduced from the X-ray diffraction data decreases with increasing pressure and iron content, in a agreement with model bases on low-pressure single-crystal elasticity data (Jackson and Khanna 1990; Sinogeikin and Bass 2000; Jacobsen et al. 2002, 2004) extrapolated using Eulerian finite strain equations (Davies 1974). This indicates that the radial diffraction data could be used to extract trends of variation of anisotropies with composition.

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